Chapter 7. Non-linear wave phenomena in GFD

V. Zeitlin

Cours GFD M2 MOCIS

Geophysical Fluid Dynamics

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Weak non-linearity vs weak dispersion: solitons

KdV equation, solitons Examples of solitons in the GFD simulations

Dispersive waves; weak non-linearity.

Non-linear dynamics of weakly non-linear Rossby waves

Dispersion relations of GFD waves and existence of resonant triads

Essentially non-linear waves

Essentially non-linear Rossby waves: modons Strongly non-linear internal gravity waves : Long

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Archetype model of breaking: "simple wave"

Non-dispersive unidimensionalwave with advective non-linearity :

$$u_t + \epsilon u u_x + c_0 u_x = 0. \tag{1}$$

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No dispersion \leftrightarrow phase velocity c_0 constant \Rightarrow breaking and shock formation.

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Unidirectional waves with weak dispersion

Dispersion of long waves:

Phase velocity: c = c(k), k - wavenumber $\rightarrow 0$. Strictly non-dispersive waves: $c = c_0 = \text{const.}$ Weak dispersion: $c = c_0 + c_1k + c_2k^2 + \dots$ Uni-directional waves: $\omega = kc(k)$ - odd function $\leftrightarrow c$ even function:

$$\omega = k(c_0 + c_2k^2 + \dots).$$

Phase-space vs physical space: Translation rules for linear systems:

$$u(k,\omega) \rightarrow u(x,t) \Rightarrow k \rightarrow -i\partial_x, \quad \omega \rightarrow i\partial_t$$

$$\omega = k(c_0 + c_2k^2 + \dots) \leftrightarrow i\partial_t u = -c_0i\partial_x u + c_2i\partial_{xxx}^3 + \dots$$

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Weak advective non-linearity + weak dispersion:

$$\partial_t u + c_0 \partial_x u + \alpha u \partial_x u + \beta \partial_{xxx}^3 u = 0.$$

Korteweg - de Vries (KdV) equation

Remark:

 c_0 can be eliminated by a change of reference frame.

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Solitary waves (solitons)

Solution propagating without change of form:

$$u = u(x - Vt) \Rightarrow$$
$$(c_0 - V)u' + \alpha uu' + \beta u''' = 0$$

Integration (localised solution - solitary wave, soliton):

$$(c_0 - V)u + \alpha \frac{u^2}{2} + \beta u'' = 0$$
(4)

Multiplication by u' and one more integration:

$$(c_0 - V)\frac{u^2}{2} + \alpha \frac{u^3}{6} + \beta \frac{u^2}{2} = 0$$
 (5)

Solution
$$u(x - Vt) = \frac{3}{\alpha} \frac{V - c_0}{\cosh^2 \sqrt{\frac{V - c_0}{4\beta}}(x - Vt)}$$
. $V = c_0 + \frac{\alpha}{3} u_{max}$

- speed depends on amplitude

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Multi-soliton solutions

Standard normalisation of the KdV equation:

$$\partial_t u + 6u\partial_x u + \partial_{xxx}^3 u = 0.$$

Solutions: $u_N = 2\partial_x^2 F_N$, *N* - number of solitons.

$$\begin{array}{rcl} F_{1} & = & 1 + e^{\eta_{1}}, \\ F_{2} & = & 1 + e^{\eta_{1}} + e^{\eta_{2}} + e^{\eta_{1} + \eta_{2} + A_{12}}, \\ F_{3} & = & 1 + e^{\eta_{1}} + e^{\eta_{2}} + e^{\eta_{3}} + e^{\eta_{1} + \eta_{2} + A_{12}}, \\ & + & e^{\eta_{1} + \eta_{3} + A_{13}} + e^{\eta_{2} + \eta_{3} + A_{23}} \\ & + & e^{\eta_{1} + \eta_{2} + \eta_{3} + A_{12} + A_{23} + A_{13}}, \\ F_{N} & = & \dots \end{array}$$

where $\eta_i = k_i x - k_i^3 t - \eta_i^{(0)}$, $A_{ij} = \left(\frac{k_i - k_j}{k_i + k_j}\right)^2$. Arbitrary initial perturbation \rightarrow series of solitons (complete integrability). Geophysical Fluid Dynamics

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Exercise: Check the formula for the bi-soliton solution *F*2.

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Solitons of equatorial Rossby waves



t=50.0 > 0 -5 20 60 'n t=100.0 80 > 0 -5 ²⁰ t=150⁴⁰.0 o 60 80 1.4 > 0 1.2 'n 20 40 60 80

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Solitons of trapped topographic waves



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Non-linear Rossby waves in the QG RSW model on the β -plane

Non-dimensional equations of motion :

$$\nabla^2 \psi_t - \psi_t + \epsilon \mathcal{J}(\psi, \nabla^2 \psi) + \psi_{\mathsf{X}} = \mathbf{0},$$

 ϵ - non-linearity parametre, $\epsilon \rightarrow 0$. Asymptotic expansion in non-linearity parametre. Solution - asymptotic series:

$$\psi = \psi^{(0)} + \epsilon \psi^{(1)} + \dots$$

Order zero: linear Rossby waves.

$$\nabla^{2}\psi_{t}^{(0)} - \psi_{t}^{(0)} + \psi_{x}^{(0)} = 0, \Rightarrow$$

$$\psi^{(0)} = \sum_{i} A_{i} e^{i(\mathbf{k}_{i} \cdot \mathbf{x} - \omega(\mathbf{k}_{i})t)} + c.c., \quad \omega(\mathbf{k}) = -\frac{k}{k^{2} + l^{2} + 1}, \quad \mathbf{k} = (k, l).$$
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Order one: first non-linear correction :

$$abla^2 \psi_t^{(1)} - \psi_t^{(1)} + \psi_x^{(1)} = -\mathcal{J}(\psi^{(0)},
abla^2 \psi^{(0)}),$$

Term in the r.h.s.:

$$\sum_{i,j} A_i A_j \left[\left(k_i l_j - k_j l_i \right) \mathbf{k}_j^2 \right] e^{i \left[\left(\mathbf{k}_i + \mathbf{k}_j \right) \cdot \mathbf{x} - \left(\omega \left(\mathbf{k}_i \right) + \omega \left(\mathbf{k}_j \right) \right) t \right]}$$
(10)
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$$\sum_{i,j} A_i A_j^* \left[\left(k_i l_j - k_j l_i \right) \mathbf{k}_j^2 \right] e^{i \left[\left(\mathbf{k}_i - \mathbf{k}_j \right) \cdot \mathbf{x} - \left(\omega \left(\mathbf{k}_i \right) - \omega \left(\mathbf{k}_j \right) \right) t \right]} + c.c.$$

Integrability conditions: solution $\psi^{(1)}$ should be bounded

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Integrability conditions:

$$\forall \hat{\psi} : \nabla^2 \hat{\psi}_t - \hat{\psi}_t + \hat{\psi}_x = \mathbf{0},$$

$$\int_{-\infty}^{\infty} dt \int_{-\infty}^{\infty} dx \int_{-\infty}^{\infty} dy \hat{\psi}^* \left(\nabla^2 \psi_t^{(1)} - \psi_t^{(1)} + \psi_x^{(1)} \right) = \mathbf{0}.$$

Therefore:

$$\int_{-\infty}^{\infty} dt \int_{-\infty}^{\infty} dx \int_{-\infty}^{\infty} dy \hat{\psi}^* \left(\mathcal{J}(\psi, \nabla^2 \psi) \right) = 0.$$
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- orthogonality of the r.h.s. to the eigen-vectors of the zero-order linear operator.

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Resonances

Necessarily: $\hat{\psi} \propto e^{i(\hat{\mathbf{k}} \cdot \mathbf{x} - \omega(\hat{\mathbf{k}})t)}$, and (11) becomes:

$$\sum_{i,j} A_i A_j \left[\left(k_i l_j - k_j l_i \right) \mathbf{k}_j^2 \right] \cdot$$

$$\int_{-\infty}^{\infty} dt \, dx \, dy e^{i \left[(\mathbf{k}_i + \mathbf{k}_j - \hat{\mathbf{k}}) \cdot \mathbf{x} - (\omega(\mathbf{k}_i) + \omega(\mathbf{k}_j) - \omega(\hat{\mathbf{k}}))t \right]}$$

$$\sum_{i,j} A_i A_j^* \left\lfloor \left(k_i l_j - k_j l_i \right) \mathbf{k}_j^2 \right\rfloor \cdot$$

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$$\int_{-\infty}^{\infty} dt \, dx \, dy e^{i \left[(\mathbf{k}_i - \mathbf{k}_j - \hat{\mathbf{k}}) \cdot \mathbf{x} - (\omega(\mathbf{k}_i) - \omega(\mathbf{k}_j) - \omega(\hat{\mathbf{k}}))t \right]} + c.c. = 0$$

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Calculus of generalised functions:

$$\int_{-\infty}^{\infty} dx e^{ikx} = \delta(k) - \text{Dirac's delta-function.}$$

Generalisation of

$$\int_0^{2\pi} dx e^{ikx} = \delta_{k0} - \text{tensor delta of Kronecker}$$

for periodic boundary conditions.

Resonances:

Non-zero contributions:

$$\mathbf{k}_{i} \pm \mathbf{k}_{j} = \hat{\mathbf{k}}, \quad \omega(\mathbf{k}_{i}) \pm \omega(\mathbf{k}_{j}) = \omega(\hat{\mathbf{k}}).$$
 (14)

three-wave resonances, resonant triads.

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Elimination of resonances

If $\exists \hat{\mathbf{k}}$ which verifies (14) the first non-linear correction is not bounded \Rightarrow asymptotic procedure is not self-consistent: resonances should be "killed".

Introducing slow evolution of the amplitudes:

$$\partial_t \to \partial_t + \epsilon \partial_T \quad \Rightarrow \tag{15}$$

$$\nabla^2 \psi_t^{(1)} - \psi_t^{(1)} + \psi_x^{(1)} = -\nabla^2 \psi_T^{(0)} - \psi_T^{(0)} - \mathcal{J}(\psi^{(0)}, \nabla^2 \psi^{(0)})$$

New contribution in the r.h.s.:

$$\sum_{i} A_{i_{\tau}} e^{i(\mathbf{k}_{i} \cdot \mathbf{x} - \omega(\mathbf{k}_{i})t)} + c.c. \Rightarrow$$
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Possibility of compensation of resonant contributions by slow evolution of amplitudes.

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A resonant triad:

$$\mathbf{k}_{1} + \mathbf{k}_{2} = \mathbf{k}_{3}, \quad \omega(\mathbf{k}_{1}) + \omega(\mathbf{k}_{2}) = \omega(\mathbf{k}_{3}), \quad (17)$$

$$\dot{A}_{3} = c(\mathbf{k}_{1}, \mathbf{k}_{2})A_{1}A_{2},$$

$$\dot{A}_{2} = c(\mathbf{k}_{3}, -\mathbf{k}_{1})A_{3}A_{1}^{*},$$

$$\dot{A}_{1} = c(\mathbf{k}_{3}, -\mathbf{k}_{2})A_{3}A_{2}^{*}, \quad (18)$$

where $c(\mathbf{k}_1, \mathbf{k}_2) = \hat{\mathbf{z}} \cdot (\mathbf{k}_1 \wedge \mathbf{k}_2)\mathbf{k}_2^2$ - interaction coefficients . This is an integrable system (in elliptic fonctions). Energy is conserved and redistributed among three waves.



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Remark:

If resonances of three waves are not possible \rightarrow resonant quartets etc.

Wave turbulence:

Ensemble of waves with random phases \Rightarrow Gaussian statistics \Rightarrow kinetic equation for the wave density, entierely determined by resonant triads (quartets) \Rightarrow energy spectra.

Successful applications

- spectra of the surface wind waves
- spectra of the internal waves in the ocean (Garret -Munk)

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Rossby waves - strong dispersion



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gravity waves : Long solutions

Barotropic inertia - gravity waves - weak dispersion



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Internal gravity waves - strong dispersion



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Demonstrate that resonant triads

- of Rossby waves in the QG RSW model on the beta-plane exist
- of inertia-gravity waves in the RSW model on the f-plane do not exist

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Essentially non-linear Rossby waves Equations of motion

$$\nabla^2 \psi_t - \psi_t + \mathcal{J}(\psi, \nabla^2 \psi) + \psi_x = 0, \qquad (19)$$

Solutions - waves propagating without change of form:

$$\psi(\mathbf{x},\mathbf{y},t) = \psi(\mathbf{x} - Ut,\mathbf{y}), \Rightarrow \mathcal{J}(\psi + U\mathbf{y}, \nabla^2\psi + (1+U)\mathbf{y}) = \mathbf{0} \Rightarrow$$

$$\nabla^2 \psi + (1+U)y = F(\psi + Uy), \qquad (20)$$

F - arbitrary function. Physical meaning: potential vorticity is constant along the streamlines.

Remark:

F is not neccessarily the same over the whole (x, y) plane. Domains with different $F \Rightarrow$ matching.

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Essentially non-linear Rossby waves: modons

Localised solutions:

$$r = \sqrt{x^2 + y^2} \to \infty \Rightarrow \psi \to 0$$

 \Rightarrow "external" *F* - linear function:

$$F(\psi + Uy) = p^2(\psi + Uy), \ p^2 = \frac{1+U}{U} > 0.$$
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Essentially non-linear Rossby waves: modons

Modons

Solutions with *F* linear outside and inside the circle r = a:

$$\nabla^2 \psi = p^2 (\psi + Uy) - (1 + U)y, \ r > a,$$

$$\nabla^2 \psi = -k^2 (\psi + Uy) - (1 + U)y, \ r < a, \quad (23)$$

Solutions in polar coordinates $x = r \cos \phi$, $y = r \sin \phi \rightarrow$ Bessel functions:

$$\psi = BK_1(pr)\sin\phi, r > a,$$

$$\psi = \left[AJ_1(kr) - \frac{r}{k^2}(1 + U + Uk^2)\right]\sin\phi, r < a (24)$$

where J_1 - Bessel function (oscillating), K_1 - modified Bessel function (decaying), A, B - constants to be determined from the conditions of matching and b.c.. Geophysical Fluid Dynamics

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Matching

Continuity of ψ , $\partial_r \psi$, $\partial_{rr}^2 \psi$:

$$\psi + Uy|_{a-} = \psi + Uy|_{a+} = 0,$$

$$\partial_r \psi|_{a-} = \partial_r \psi|_{a+}.$$

2 first conditions giving A, B:

$$A=\frac{a(1+U)}{k^2J_1(ka)}, \quad B=-\frac{Ua}{K_1(pa)}.$$

3rd condition: determining k:

$$\frac{J_1'(ka)}{J_1(ka)} = \frac{1}{ka} \left(1 + \frac{k^2}{p^2} \right) - \frac{k}{p} \frac{K_1'(pa)}{K_1(pa)}.$$
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 \forall (*a*, *p*) infinite series of solutions for *k*. Minimal value of *k* - dipolar structure.

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Streamfunction of a modon



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Strongly non-linear internal gravity waves : Long solutions

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Exercice:

Obtain the formulas (26), (27).

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Modons and atmospheric blockings



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Stratified non-rotating fluid

2d Boussinesq equations:

$$u_t + uu_x + ww_z + \phi_x = 0,$$

$$w_t + uw_x + ww_z + \Xi + \phi_z = 0,$$

$$u_x + w_z = 0, \quad \Xi_t + u\Xi_x + w\Xi_z = 0.$$
 (28)

 $\Xi = \frac{g(\rho(z)+\sigma)}{\rho_0}$ - buoyancy variable, including the effects of background stratification $\rho(z)$, $\phi = \frac{P}{\rho_0}$ - geopotential, ρ_0 - constant normalisation density, σ - density perturbation.

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Equations in streamfunction - buoyancy variables

$$\Delta \psi_t + \mathcal{J}(\psi, \Delta \psi) + \Xi_x = 0,$$

$$\Xi_t + \mathcal{J}(\psi, \Xi) = 0.$$
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 ψ - streamfunction, $\zeta=-\Delta\psi$ - horizontal vorticity, Δ - Laplacian, ${\cal J}$ - Jacobian.

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Weak non-linearity vs weak dispersion: solitons

KdV equation, solitons Examples of solitons in the GFD simulations

Dispersive waves; weak non-linearity.

Non-linear dynamics of weakly non-linear Rossby waves

Dispersion relations of GFD waves and existence of resonant triads

Essentially non-linear waves

Essentially non-linear Rossby waves: modons

Hydrostatic limit (long waves)

Replacement of the equation for *w* by hydrostatic equation $-\Xi = \phi_z$:

$$\psi_{zzt} + \mathcal{J}(\psi, \psi_{zz}) + \Xi_x = \mathbf{0},$$

$$\Xi_t + \mathcal{J}(\psi, \Xi) = \mathbf{0}.$$

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(30)

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Strongly non-linear internal gravity waves : Long solutions

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Stationary solutions (change of reference frame \Rightarrow propagation at constant speed):

$$\begin{aligned} \mathcal{J}(\psi,\Delta\psi) + \Xi_{\mathbf{X}} &= \mathbf{0}, \\ \mathcal{J}(\psi,\Xi) &= \mathbf{0}. \end{aligned}$$

Therefore $\Xi = \Xi(\psi)$ and

$$\begin{aligned} \mathcal{J}(\psi,\Delta\psi) + \Xi'(\psi)\psi_{\mathsf{X}} &= 0 \quad \Rightarrow \\ \mathcal{J}(\psi,\Delta\psi + \Xi'(\psi)z) &= 0 \quad \Rightarrow \\ \Delta\psi + \Xi'(\psi)z &= F(\psi), \end{aligned}$$

(32)

(31)

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where $\Xi(\psi)$ and $F(\psi)$ - arbitrary functions.

Long's waves

Upstream ($x \to \infty$): b.c. of constant velocity $\to \psi = cz$, and of a given stratification $\Xi = \Xi_0(z) \Rightarrow$

$$\Xi(\psi) = \Xi_0(\frac{\psi}{c}), \quad F(\psi) = \Xi'(\psi)\frac{\psi}{c} = \frac{\psi}{c^2}\Xi'_0(\frac{\psi}{c})$$
(33)

Example: linear stratification upstream: $\Xi_0 = \text{const} + \alpha z$ \rightarrow Long's equation (linear!) for a non-linear stationary wave:

$$\Delta \psi + rac{lpha}{c}(z - rac{\psi}{c}) = 0.$$

New variable - deviation of streamlines:

$$\phi = \psi - cz, \Rightarrow \Delta \phi - \frac{\alpha}{c^2} \phi = 0.$$
 (35)

B.c. in *z*: $\phi|_{z=h(x)} = h(x)$, *h* - topography.

(34)

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Exercice:

- Consider hydrostatic version of (34). Find its particuliar solution verifyi ng automaticaly the b. c. at z = h(x).
- In a half-bounded domain in z, what should be a b. c. at z → ∞ to be used?

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